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# Characterization of the field saturated hydraulic conductivity on a hillslope: measurement techniques, data sensitivity analysis and spatial correlation modelling

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## Abstract

In the context of studies aiming at the estimation of effective parameters for unsaturated zone modelling, this work tackles the problem of experimental data quality, considering the large collection of data gathered at an experimental site equipped for unsaturated zone hydraulic monitoring in the alluvial basin of a Calabrian river, in the South of Italy.

Focusing attention on field saturated hydraulic conductivity, the in-site measurement techniques by tension disc and pressure ring infiltrometers are considered, pointing out the main indications for the correct use of each measuring approach; laboratory techniques are also considered.

Statistical data analysis showed that the measurements performed by tension disc infiltrometer supplied values of hydraulic conductivity which are on average lower and more homogeneous than the values provided by the other measurement techniques considered.

Sensitivity analysis was then carried out by Monte Carlo simulation on the parameter sampling achieved by field measurement techniques in order to evaluate the influence of any possible small measurement errors on the data.

Sensitivity analysis showed that both ring and disc infiltrometer are tools reliable enough for the in situ measurements of field saturated hydraulic conductivity.

Finally, after a data merging procedure giving origin to different sets of data, the spatial correlation structure of field saturated hydraulic conductivity is investigated, using well-known geostatistical techniques.

## 1. Introduction

Research about water flow and pollutant transport processes in an unsaturated zone is one of the topic goals for hydro-geologists and soil scientists, because environmental protection and recovery depends on the possibility of forecasting the hydraulic behaviour of the subsoil.

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Subsoil system modelling is currently supported by physically based mathematical models. The most reliable models in hydrogeology, among the different types of applicable models, are those characterized by spatially distributed parameters, because they are able to take into account the complex geometry and physical heterogeneity of subsurface natural systems.

The practical use of this type of model, however, requires experimental measurements and the characterization of hydraulic properties, such as hydraulic conductivity, on the numerical mesh representing the subsurface system under analysis.

In unsaturated subsoil, hydraulic conductivity is a non-linear function of the matrix potential and water content. In the opinion of many researches, the characterization of unsaturated zone based on the use of experimental acquisition of hydraulic conductivity data is a time-consuming and costly task, while data collection about water content and pressure head seems easier and quicker. Therefore, an inverse problem arose: how to achieve subsurface hydraulic parameter characterization by experimental sampling prevalently of dependent variables, that is, soil moisture and matrix potential, rather than of hydraulic conductivity.

The modern approach to the inverse problem is based on the concept of effective parameters, which points out the problem of reducing the gap between the measurement scale and the model scale, in order to obtain the best possible reliable modelling issues.

To estimate the effective values of hydraulic conductivity two approaches were proposed in the scientific literature: the former based on up-scaling techniques (Bierkens et al., 2000; Khaleel et al., 2002), the latter based on optimisation techniques (Yeh and Zhang, 1996; Zhang and Yeh, 1997; Gupta et al., 1998).

In the context of studies aiming at effective parameter estimation by the optimisation approach, this paper tackles the problem of experimental data quality, referring to the large collection of water content, matrix potential, and hydraulic conductivity data, gathered at an experimental site equipped for unsaturated zone hydraulic monitoring in the alluvial basin of a Calabrian river, in the South of Italy. In particular, focusing attention

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on field saturated hydraulic conductivity  $K_s$ , after some reminders about the measurement techniques used (ring infiltrometer, disc infiltrometer, laboratory tests), this work illustrates the results of data statistical analysis carried out in order to investigate the effects of the different techniques of measurement on the geovariance of  $K_s$ . Sensitivity analysis was then carried out by Monte Carlo simulation on disc and ring infiltrometer data in order to evaluate the influence of eventual small measurement errors on the data obtained.

Finally, after a data-merging procedure giving origin to different sets of data, the spatial correlation structure of  $K_s$  is investigated, using well-known geostatistical techniques.

**2. Methods**

2.1. Water movement in the unsaturated zone

In unsaturated porous media water flow is described by Richards’s equation, derived by combining the Darcy-Buckingham equation with mass balance principle. The 1-dimensional Richards’s equation is written as follows:

$$C(h) \frac{\partial h}{\partial t} = \frac{\partial}{\partial z} \left[ k(h) \cdot \frac{\partial h}{\partial z} \right] + \frac{\partial k(h)}{\partial z} \tag{1}$$

with

$$C(h) = \frac{\partial \vartheta}{\partial h}, \tag{2}$$

where  $h[L]$  is the soil water head,  $\vartheta [L^3 L^{-3}]$  is the volumetric water content,  $z[L]$  is the distance of water vertical path,  $t[T]$  is time,  $K(h) [LT^{-1}]$  is the unsaturated hydraulic conductivity,  $C(h)$  is the soil water capacity  $[L^{-1}]$  (Richards, 1931).

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The hydraulic behaviour of unsaturated soils is characterized by soil hydraulic functions such as: the soil water retention curve  $\vartheta(h)$ , and the hydraulic conductivity curve  $K(h)$  or  $K(\vartheta)$ .

2.2. Measurement techniques

5 The hydraulic properties of unsaturated soil can be experimentally defined both by in situ or laboratory measurement methods.

Among the several field techniques used to measure the hydraulic properties of soils in the unsaturated zone, infiltration-based methods showed viability and effectiveness more than others. In particular, with respect to the data of concern in this work, two tools were used: a single ring pressure infiltrometer (Vauclin et al., 1994; Mertens et al., 2002) and a tension disc infiltrometer (Perroux and White, 1988). Both tools allow the estimation of field saturated hydraulic conductivity  $K_s$ . Because of entrapped air during the infiltration process, field saturated hydraulic conductivity could be lower than hydraulic conductivity in completely saturated conditions. Thereby,  $K_s$  will be referred to as  $K_{fs}$  or  $K_{ts}$  respectively for ring infiltrometer or disc infiltrometer test.

2.2.1. Single ring pressure infiltrometer measurements

The ring infiltrometer allows the estimation of the matric flux potential  $\phi_m$ :

$$\phi_m = \int_{h_i}^0 K(h)dh, \tag{3}$$

where  $h_i \leq 0$  is the initial value of soil water head (Elrick and Reynolds, 1989).

20 As consequence, it is possible also to estimate the parameter  $\alpha_m$ , given by:

$$\alpha_m = \frac{K_{fs}}{\phi_m}, \tag{4}$$

which provides a measure of the relative contributions of gravity and capillary forces to water movement in an unsaturated soil.

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According to steady-flow conditions, flow rate out of a ring infiltrometer is given by:

$$Q_{PI} = K_{fs} + \frac{H}{\pi \cdot r_d G} K_{fs} + \frac{\phi_m}{\pi \cdot r_d G}, \quad (5)$$

where  $H[L]$  is the steady value of hydraulic head applied during the measurement,  $G$  is a shape parameter given by

$$G = 0,316 \frac{d_r}{r_d} + 0,184 \quad (6)$$

and  $r_d [L]$  is the ring radius,  $d_r$  is the depth of insertion of the ring into the soil (Reynolds and Elrick, 1990).

According to the Simultaneous Equation Approach (S.E.A.), two values of head ( $H_1$  and  $H_2$ ) were applied successively, in order to obtain the system of two equations from which the two unknown parameters  $K_{fs}$  and  $\phi_m$  (or  $\alpha_m$ ) can be determined.

For an effective interpretation of ring infiltrometer measurements, some warnings should be considered. In the first place, it is important to underline that  $K_s$  and  $\alpha_m$  are calculated by the observed infiltration rates ( $Q_1$  and  $Q_2 [LT^{-1}]$ ), which correspond to the applied heads ( $H_1$  and  $H_2$ ), and that  $K_s$  and  $\alpha_m$  are strongly dependent on the ratio  $Q_2/Q_1$  (Philip, 1985).

$Q_2/Q_1$  ratio can be affected by different source of error: a) Non-attainment of true steady state flow; b) Error in the experimental measurements of  $Q_1$  and  $Q_2$  because of air bubble size and reading errors in the permeameter; c) Entrapment of air in the soil due to water redistribution during filling of the permeameters and restarting of infiltration (Heinen and Raats, 1990; Wu et al., 1993; Russo et al., 1997; Elrick and Reynolds, 1992).

The accuracy of measurements by single ring pressure infiltrometer can be affected by other factors: during the insertion of the infiltrometer ring, soil macrostructure can collapse for soil matrix compression (or soil shattering), and the wall of the ring can cut off non-vertical pores.

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### 2.2.2. Single disk tension infiltrometer measurements

For the tension disk infiltrometer the tension flux potential ( $\phi_t$ ) is given by:

$$\phi_t = \int_{\psi_i}^{\psi_t} K(\psi) d\psi \quad (7)$$

with

$$5 \quad \psi_i \ll \psi_t \leq 0, \quad (8)$$

where  $\psi_t$  is the constant water potential applied on the soil surface (Elrick and Reynolds, 1992).

Likewise for the ring infiltrometer, it is possible to define the parameter  $\alpha_{mt}$ , given by:

$$\alpha_t = \frac{K_{ts}}{\phi_t} \quad (9)$$

10 The interpretation of disc infiltrometer measurements is also based on an S.E.A.; according to the Wooding algebraic approximation of steady-state unconfined infiltration rate into soil from a circular source of radius  $r$  (Wooding, 1968), the following relation was used (Angulo-Jaramillo et al., 2000):

$$Q_{TI} = \pi \cdot r^2 K_{ts} e^{\alpha \psi_t} \left[ 1 + \frac{4}{\pi \cdot r \alpha_t} \right] \quad (10)$$

15 Also for disc infiltrometer measurement interpretation some warnings should be taken in account.  $K_{ts}$ -values obtained by such a measuring device can be affected by errors and the accuracy of data depends on several critical factors (Elrick and Reynolds, 1992; Ankeny et al., 1991).

### 2.2.3. Laboratory measurements

Regarding the laboratory measurements method, having taken a number of undisturbed soil samples, the saturated hydraulic conductivity ( $K_s$ ) for each of them can be easily determined using a constant head permeameter (Klute, 1986). Of course, the most critical constraint of laboratory measurement procedures is the condition of undisturbed sampling, which actually is very difficult to achieve.

### 2.3. Statistical analysis

Hereinafter the symbol  $K_s$  will be used to indicate both saturated hydraulic conductivity  $K_{fs}$  and  $K_{ts}$ , measured respectively for ring infiltrometer or disc infiltrometer test; in the same manner the symbols  $\phi_m$  and  $\alpha_m$  will be used to indicate respectively the parameters  $\phi_m$ ,  $\phi_t$  and  $\alpha_m$ ,  $\alpha_t$ .

Descriptive conventional statistics of  $K_s$  based on the available data are achieved in order to compare the results provided by the different measurement techniques applied. Univariate distribution can be tested, by calculating relative frequency histogram and cumulative probability distribution. A comparison between  $K_s$ -values obtained in the same locations by different measurement techniques can be a help to focus the advantages and limits of each measuring approach.

### 2.4. Monte Carlo based sensitivity analysis

Based on Monte Carlo techniques of numerical random generation it is possible to carry out a simple but effective error analysis of  $K_s$  and  $\alpha_m$  data collected by using ring and disc infiltrometers. The aim of such an error analysis is to assess the propagation of small random error in evaluating the steady-state inflow rates on the data acquisition of  $K_s$  and  $\alpha_m$ . The error is assumed equivalent to a possible reading-off error on the reservoir of the measurement tools. The hypothesis of a normal distribution characterized by mean 0 mm and standard deviation equal to 1 mm is assumed. A number

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of 1000 Monte Carlo simulations of reading errors is performed for each measurement location. The error is then added to read-off values and  $K_s$  and  $\alpha_m$  re-calculated using the SEA for each random error. Finally, considered the mean value  $\mu$  and standard deviation  $\sigma$  of  $K_s$  and  $\alpha_m$  in each measurement location, a diagram  $\mu-\sigma$  both for  $K_s$  and  $\alpha_m$ . The trend and scattering of standard deviation can be therefore evaluated in order to obtain conclusions about  $K_s$  and  $\alpha_m$  sensitivity to small measurement errors (Saltelli et al., 2000).

2.5. Geostatistical analysis

The spatial correlation structure of  $K_s$  in the unsaturated zone is investigated applying geostatistical techniques such as the calculation based on the available data of a few points of the semivariogram function, which, with respect to a generic variable  $Z$  spatially distributed, is given by

$$\gamma(h_l) = \frac{1}{2n(h_l)} \sum_{i=1}^{n(h_l)} [z(x_i) - z(x_i + h_l)]^2, \tag{11}$$

where  $\gamma(h_l)$  is the estimated semivariance for lag distances class  $h_l$ ,  $z(x_i)$  and  $z(x_i + h_l)$  are the measured sample values at point  $x_i$  and  $x_i + h_l$  respectively, and  $n(h_l)$  is the total number of sample pairs for the distance interval  $h_l$  (Mallants et al., 1997). A variogram model is then searched in order to fit the experimental points (Goovaerts, 1997).

### 3. Characterization of the field saturated hydraulic conductivity on a hill slope: materials

#### 3.1. The experimental site: hydrogeological settings and measurement locations

The experimental site, located in Calabria (South Italy), in the basin of the Turbolo river, the left affluent of the Crati river, is a flat alluvial area with a surface of 2800 m<sup>2</sup> (70 m×40 m) and a mean slope of about 6% (Figs. 1a and 1b). Thanks to the collaboration between the Institute for Land and Water Management of the Katholieke Universiteit of Leuven (Belgium) and the Department of Soil Conservation of the University of Calabria (Italy) it was equipped for monitoring the soil water content ( $\theta$ ) at different depths and assessing other precipitation and runoff processes; moreover, measurements of pressure heads were performed, using special devices and probes (TDR, tensiometers, gutters, etc). Figure 2 shows a view the experimental site.

The aim of this work, however, is the characterization of field saturated hydraulic conductivity; therefore, hereinafter, just the  $K_s$ -data collected in a restricted zone (40 m×20 m) of the whole experimental site will be dealt with. No other mention will be made to the other data collected.

Geological settings of the area include alluvial deposits with prevalence of sand and conglomerates and a widespread presence of loam. The shallow aquifer is located at 8.3 m from the ground surface with a thickness of about 8 m.

Soon after the summer of 2002  $K_s$ -measurements began, 30 of which were carried out by single ring infiltrometer and 30 by disc infiltrometer. In addition 30 undisturbed soil samples were taken in the same locations of single ring infiltrometer measurement. Figure 3 shows in situ measurement locations and undisturbed soil sampling points.

The choice of the measurement period was not casual. In fact, Eqs. (3) and (7) show that the flux potential is a function of the initial soil water head. Then the measurement process of  $K_s$  could be affected by the initial moisture conditions of the site. Dependence of flux potential on the initial soil water head, however, can be neglected if the initial soil water head is very negative, that is, in other words, if the initial soil is dry

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(Bagarello and Provenza, 1996; Russo et al., 1997).

For this reason, the single ring pressure infiltrometer measurements were carried out after a long dry period. Five data of  $K_s$  and  $\alpha_m$  obtained by single ring pressure infiltrometer were discarded because they showed negative values. In the same manner, two anomalous values determined in laboratory were cut off from the data set. On the contrary, no negative values was provided by the tension disc infiltrometer measurements.

## 3.2. Results and discussion

### 3.2.1. Statistical analysis of data

In order from a statistical point of view to characterize each data set, derived by using different measuring techniques, some univariate statistics, such as mean value, standard deviation and asymmetry, were considered; in particular the asymmetry was calculated by the Pearson coefficient. Table 1 and Fig. 4 summarize and compare statistics obtained for each data source. A log scale needs to be used to keep together the resulting variety of magnitude orders in the same diagram (Jury, 1985; Zavattaro et al., 1999; Mohanty et al., 1994; Jacques, 2000). Figures 5a and 5b, 6a and 6b, 7a and 7b, report the relative frequency histogram for each data set, both on the natural and on the logarithmic scale. Figures 5c and 5d, 6c and 6d, 7c and 7d, report the probability plot for each data set, both on the natural and on the logarithmic scale.

From the above-mentioned figures the following elements can be observed: a) The disc infiltrometer provides data ranging from about 50 to about 600 cm/day, enough homogeneous but asymmetric with respect to the mean value; b) The ring infiltrometer provides data ranging from about 100 to about 37 000 cm/day, very spread and enough symmetric with respect to the mean value; c) The laboratory provides data ranging from about 1 to 22 000 cm/day, very spread and asymmetric with respect to the mean value.

Such a statistical comparison of  $K_s$  data obtained with different measurement tech-

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niques provides results which are coherent with the conclusions of other studies on the  $K_s$  estimate problem (Herman et al., 2003).

The logarithmic transformation of data, however, corrects the asymmetry of the disc infiltrometer data, and produces a distribution less spread and asymmetric for the ring infiltrometer data. For the lab data set, on the contrary, the asymmetry is reduced but the standard deviation remains high. In addition, lab data distribution seems tri-modal. In fact, for disc and ring infiltrometer data the asymmetry is due to few outliers, much lower or higher than the data trend. On the contrary, the lab data set seems originated by three different data sources, rather than by a trend component plus few outliers.

### 3.2.2. Comparison of ring infiltrometer and laboratory data

A specific comparison was made between ring infiltrometer and lab data, because measurements were achieved in the same locations (Reynolds et al., 2000). The comparison was possible in 22 locations. Figures 8a and 8b represent the results of such a comparison. Ring infiltrometer measurements, except for two locations, show values much higher than lab data, which, in many cases, appear very low.

Taken the two  $K_s$ -values recorded in a fixed location respectively by ring infiltrometer and laboratory measurement method, the ratio between the lowest value and the highest one provides a likelihood index ( $L_j$ ), which can be used to compare the order of magnitude of ring infiltrometer data versus the laboratory ones.

Figures 9a and 9 b report such a likelihood index  $L_j$ , pointing out that: a) ring infiltrometer data and laboratory data show the same order of magnitude just in seven cases; b) the lowest values of  $L_j$  occur prevalently for low values of lab  $K_s$ -data.

### 3.2.3. Sensitivity analysis of disc and ring infiltrometer data

Sensitivity analysis was carried out separately on ring and disc infiltrometer data by using Monte Carlo simulations to generate the random errors affecting the reading of levels recorded by the reservoirs of the measurement tools. Figures 10a and 10b

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represent the results of sensitivity analysis for ring infiltrometer measurements. Figures 11a and 11b represent the results for disc infiltrometer measurements. The scatter-diagrams in the above mentioned figures represent the standard deviation versus mean values of  $K_s$ , and  $\alpha_m$ .

The statistical elaboration of Monte Carlo simulations show that: a) Small random errors made during the test management affect the ring infiltrometer measurement of the hydraulic conductivity less than the disc infiltrometer measurement of  $K_s$ ; in fact, the standard deviation of the former is lower for an order of magnitude than the standard deviation of the latter. Both ring and disc infiltrometer measurements, however, show standard deviation distributions without specific trend and more or less accumulated around low values; b) Concerning  $\alpha_m$  measurement, standard deviation vs mean values graphics, represented on the logarithmic scale, show, on the contrary, that the ring infiltrometer is more sensitive to small random measurement errors than the disc infiltrometer. In addition, the standard deviation of  $\alpha_m$  in the case of the ring infiltrometer is characterized by an increasing drift which is absent for the disc infiltrometer measurements.

### 3.2.4. Geostatistical analysis

#### a) Data merging

Eight data sets were created by merging the disc infiltrometer  $K_s$ -measurements with variable increasing portions of  $K_s$ -data, deduced from laboratory tests or ring infiltrometer measurements, as explained hereinafter. Four data sets include disc tension infiltrometer and laboratory data; the remaining four data sets include disc tension infiltrometer and ring pressure infiltrometer data.

These data sets were built following the criteria illustrated in Table 2. The likelihood index, previously defined and represented in Figs. 9a and 9b, was considered in order to select each time which portion of laboratory or ring infiltrometer data should be added to the disc infiltrometer data set.

Therefore, each one of the four data sets labelled as “disclab” was obtained by merg-

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ing all the disc infiltrometer data with different portions of the laboratory data, which was defined according to assigned values of  $L_j$ , as shown in Table 2. It should be observed that the fourth data set is constituted of all the available data collected both by the disc infiltrometer and the laboratory test. In the same manner the four data sets labelled as “discring” were obtained by joining all the disc infiltrometer data with variable quotas of the ring infiltrometer data (see Table 2); also in this case the fourth data set contains all the available data obtained both by the disc and the ring infiltrometer. Table 2 also shows the total number of data considered for each data set.

*b) Spatial correlation structure analysis*

Using each of the above mentioned data sets, the spatial correlation structure of  $K_s$  in the experimental site considered was investigated.

Figure 12 summarizes the experimental variograms obtained for the data sets formed by the disc infiltrometer data plus increasing portions of laboratory data. Figure 13 shows the experimental variograms obtained for the data sets formed by the disc infiltrometer data plus increasing portions of the ring infiltrometer data. Figure 12 points out a non-stationary behaviour for  $K_s$ . Figure 13, on the contrary, shows that the non-stationary behaviour of  $K_s$  is confirmed just for the data set labelled as “discring1”, while for the other ones the behaviour of  $K_s$  gradually trends to become stationary, with the achievement of a well-defined sill.

Hence, as was expected, the results of spatial correlation analysis are similar if the laboratory and ring infiltrometer data with high  $L_j$  (disclab1 variogram vs discring1 variogram) are added to the disc infiltrometer data.

On the contrary, the variograms show very different spatial structure for the data sets involving the laboratory and ring infiltrometer data with low  $L_j$ .

Figures 14a, 14b and 15a, 15b show variogram models which were selected for the above-mentioned data sets. The aim was to compare the spatial structure of  $K_s$ -data according to the variation of  $L_j$ , that is, to investigate how  $K_s$ -data spatial structure changes when data with decreasing  $L_j$  are involved in the analysis.

As expected for the lowest values of  $L_j$  the spatial structure of  $K_s$ -data is very different

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if the data sets labelled as “disclab” are considered instead of the data sets labelled as “discring”.

Figures 15a and 15b show how a non-stationary model should be always assumed for the “disclab” data sets while a stationary model can be adopted for the “discring” data sets when the portion of the ring infiltrometer data increases, that is, when  $L_j$  decreases.

#### 4. Conclusions

The present paper, referring to the field saturated hydraulic conductivity measurements performed in an experimental site purposely equipped in the unsaturated zone of the Turbolo basin (Calabria, Italy), faced the problem of data quality achieved either by laboratory analysis, on undisturbed samples, or by field devices (tension disc infiltrometer and pressure ring infiltrometer).

Statistical analysis of  $K_s$  data showed that the measurements performed by the tension disc infiltrometer supplied  $K_s$ -values which are on average lower and more homogeneous than other ones; moreover, the starting asymmetry of the data set is effectively corrected by a simple logarithmic transformation. Data acquired by the pressure ring infiltrometer and those acquired in the laboratory, on the contrary, are characterized by a marked asymmetry and by a high scattering around the mean value, as well as by several outliers, much higher or much lower than the mean value.

Specifically, the laboratory data seem to be formed by three data subsets which are mutually uncorrelated, above all because of the high influence of very low hydraulic conductivity values. So much so that the logarithmic transformation, when applied to  $K_s$ -values obtained by pressure ring infiltrometer, is able to lead the behaviour of data again to a gaussian scheme, while, when applied to the laboratory data, it produces a 3-modal distribution.

Comparison of the data obtained in the same locations, by pressure ring infiltrometer and by laboratory analysis, showed that about 1/3 of the data has the same order of

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magnitude and that the highest differences occurs in correspondence to the lowest laboratory data.

Sensitivity analysis, carried out on  $K_s$ -field data, using Monte Carlo simulations, showed that the pressure ring infiltrometer is a tool reliable enough for the in situ determination of  $K_{fs}$ , while it does not produce such stable data for the parameter  $\alpha_m$ .

The measurements performed by tension disc infiltrometer show, on the contrary, a reliable enough level for the in situ estimation either of  $K_{ts}$ , or  $\alpha_m$ .

Geostatistical analysis, performed on data sets obtained joining the tension disc infiltrometer data with progressively increasing portions of the data gathered respectively by laboratory analysis and by pressure ring infiltrometer, showed that: a) the addition of laboratory data produces variograms with non-stationary behaviour, which is confirmed whatever is the entity of the data set analyzed every time; b) the addition of pressure ring infiltrometer data produces variograms which, starting from a non-stationary state, gradually behave like stationary spherical functions, characterized by a very small correlation scale (about 5÷6 m).

The results of  $K_s$ -spatial analysis, obtained by merging and analyzing only in situ sampling data, are similar to those which can be found in analogous studies referred to in the literature; on the contrary, the results obtained by analyzing data sets which were formed both by field and laboratory data show a non-stationary behaviour which cannot be ascribed to the physical presence of a drift component. Therefore, it could be due to the more irregular composition of the laboratory data.

*Acknowledgements.* The authors thank the J. Feyen, Director of the Institute for Land and Water Management of the Katholieke Universiteit Leuven (Belgium), for his precious suggestions and to have allowed this research, furnishing the measurement devices and performing the laboratory tests; furthermore they thank, also, the J. Mertens, Research Hydrologist of Lincoln Environmental Ruakura Research Centre (Hamilton, New Zealand), that participated to the initial stage of the research.

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**Table 1.** Summary of data simple statistics.

Statistics	Data (cm/s)		
	Disc infiltrometer	Ring infiltrometer	Laboratory
Asymmetry	2.5146	2.0834	3.2204
Standard deviation	0.0012	0.0996	0.0553
Median value	0.0018	0.0396	0.0008
Mean value	0.0022	0.0842	0.0236
Minimum value	0.0006	0.0015	1.743E-05
Maximum value	0.0069	0.4256	0.2546

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**Table 2.** Data set composition.

Data set label	Merging criteria	Total number of data
Disclab1	Disc infiltrometer data + laboratory data for which $L_i \geq 10^{-1}$	37
Disclab2	Disc infiltrometer data + laboratory data for which $L_i \geq 10^{-2}$	43
Disclab3	Disc infiltrometer data + laboratory data for which $L_i \geq 10^{-4}$	52
Disclab	Disc data + all laboratory data	58
Discring1	Disc data + ring infiltrometer data for which $L_i \geq 10^{-1}$	37
Discring2	Disc data + ring infiltrometer data for which $L_i \geq 10^{-2}$	43
Discring3	Disc data + ring infiltrometer data for which $L_i \geq 10^{-4}$	52
Discring	Disc data + all ring infiltrometer data	55

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Fig. 1. (a) Geographical location of Turbolo basin.

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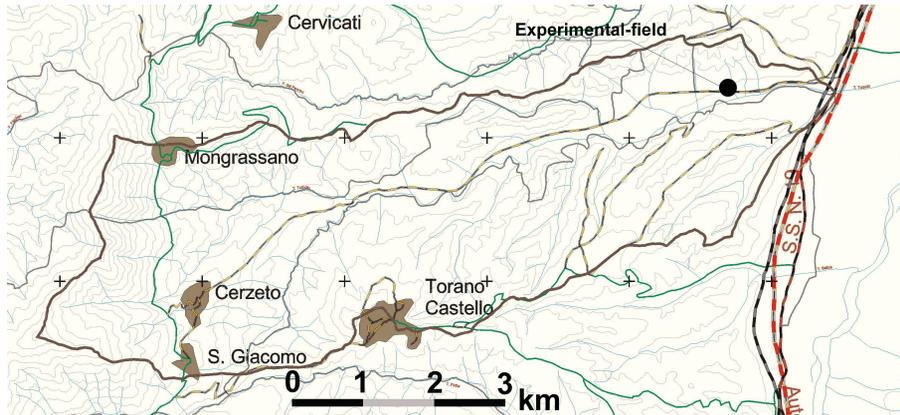


Fig. 1. (b) Turbolo basin delimitations.

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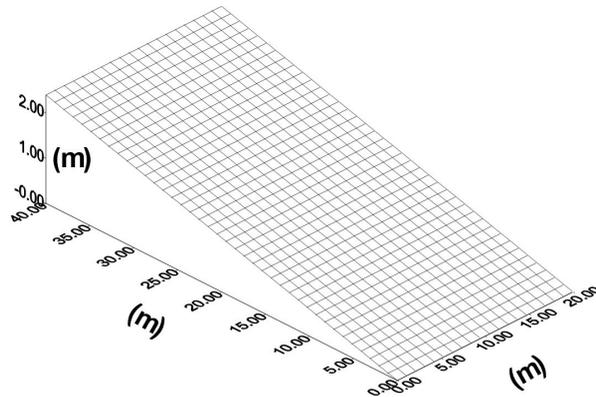
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**Fig. 2.** View of the experimental site.

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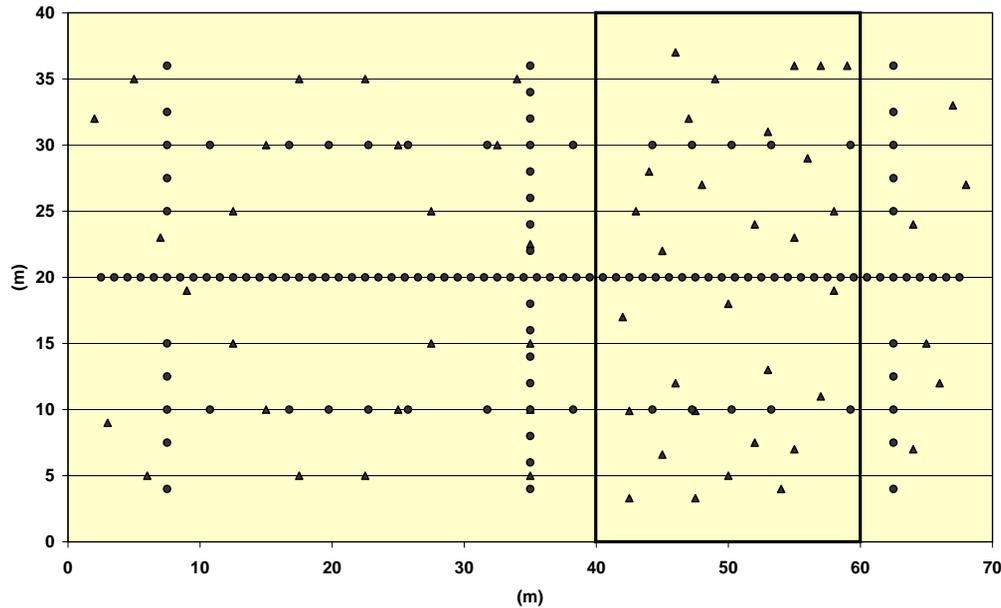
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**Fig. 3.** Measurement locations. ● Tension disc infiltrometer measurement locations. ▲ Pressure ring infiltrometer measurement locations and undisturbed soil sampling points.

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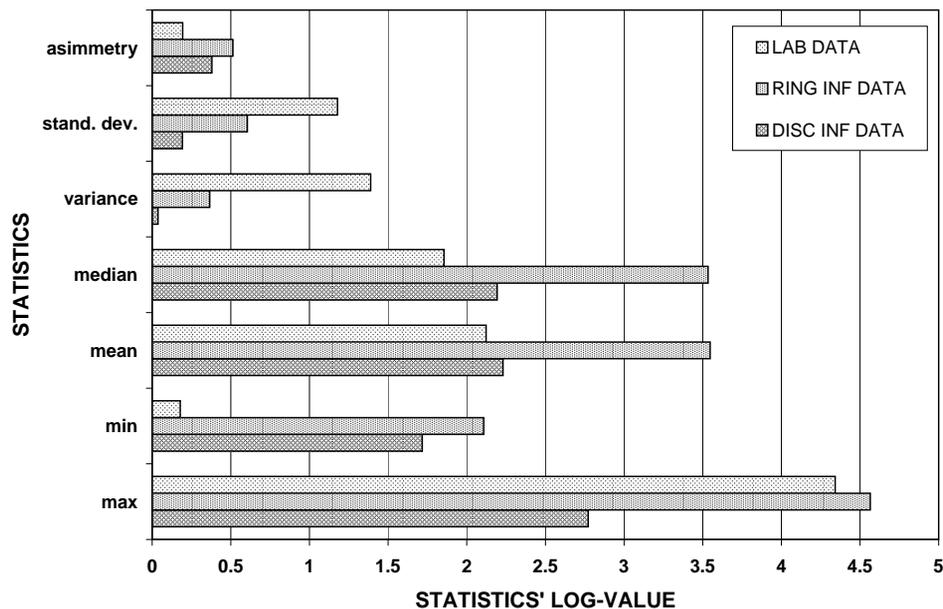


Fig. 4. Comparison of  $K_s$ -data main statistics.

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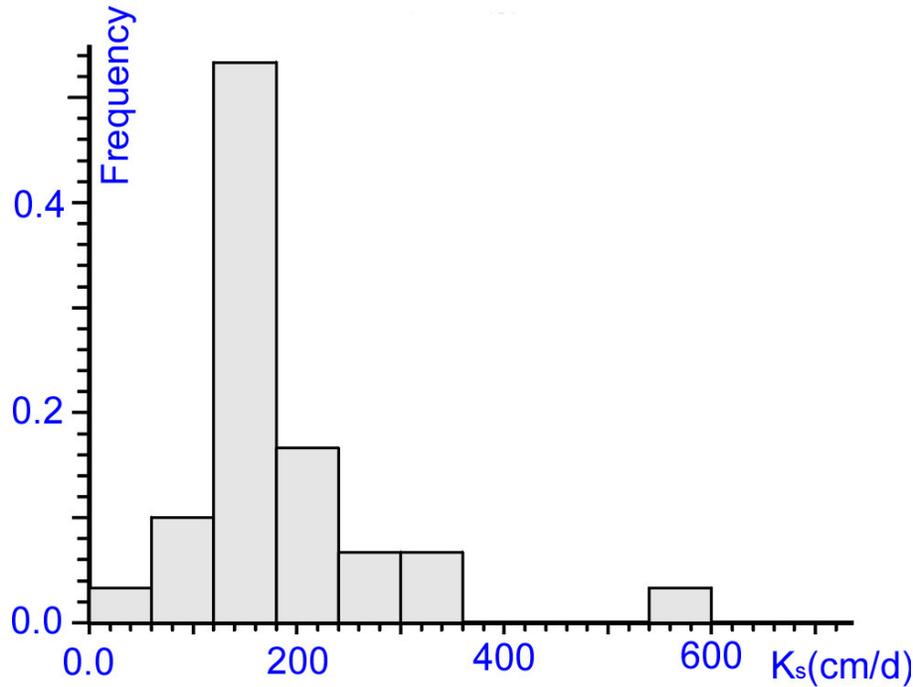
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**Fig. 5. (a)** Disc infiltrometer data histogram.

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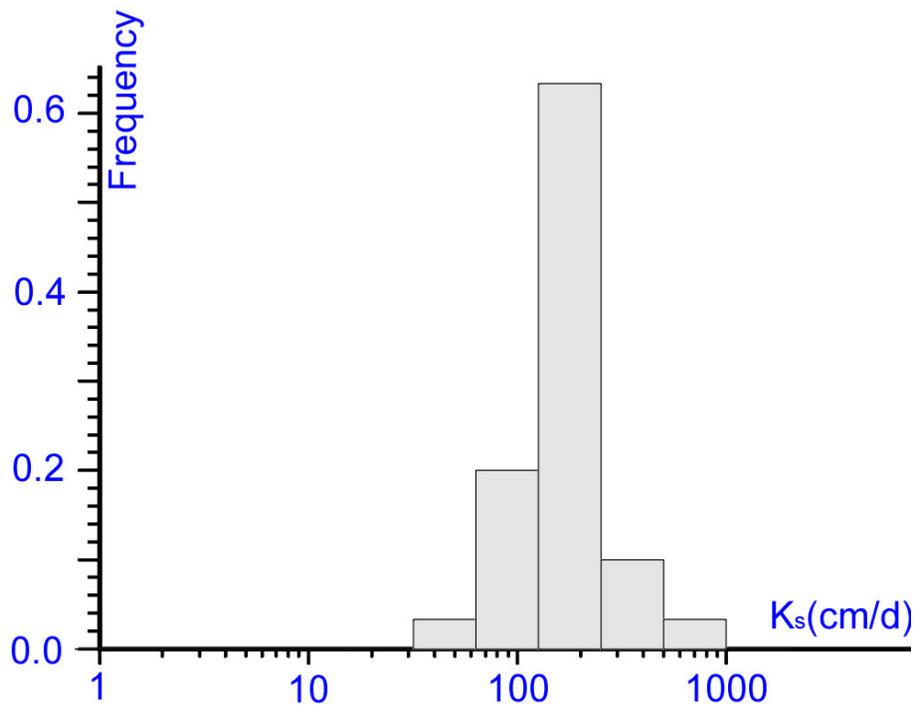


Fig. 5. (b) Disc infiltrometer data log-histogram.

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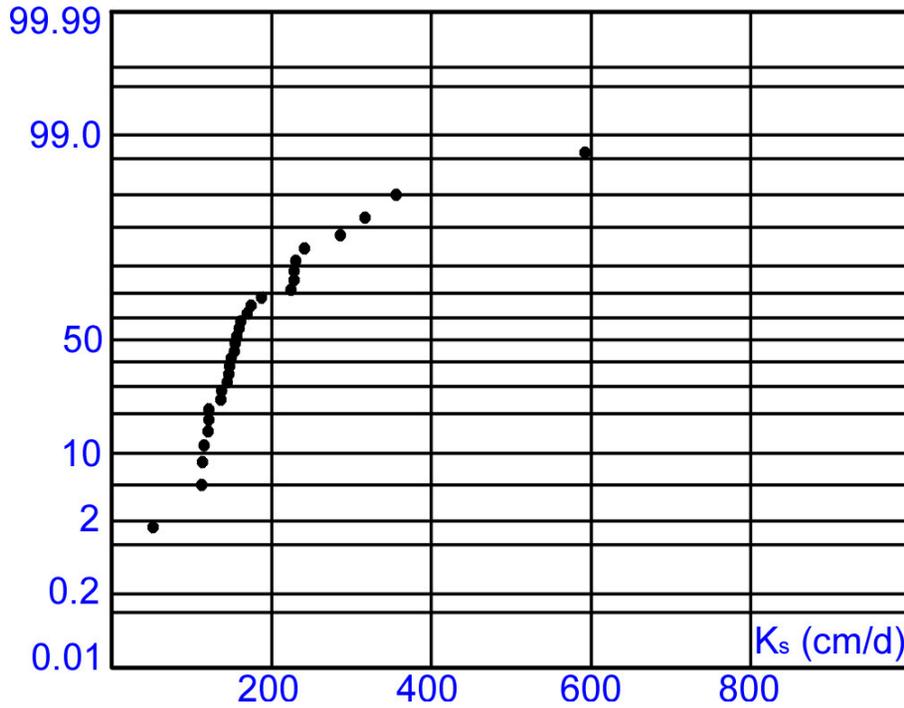


Fig. 5. (c) Disc infiltrometer data probability plot.

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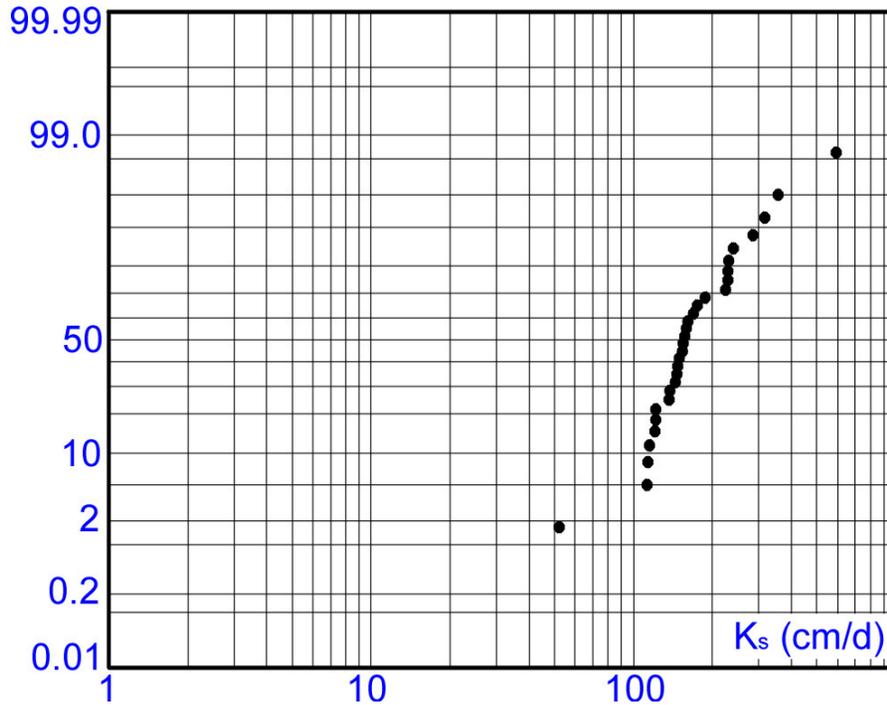


Fig. 5. (d) Disc infiltrometer data log-probability plot.

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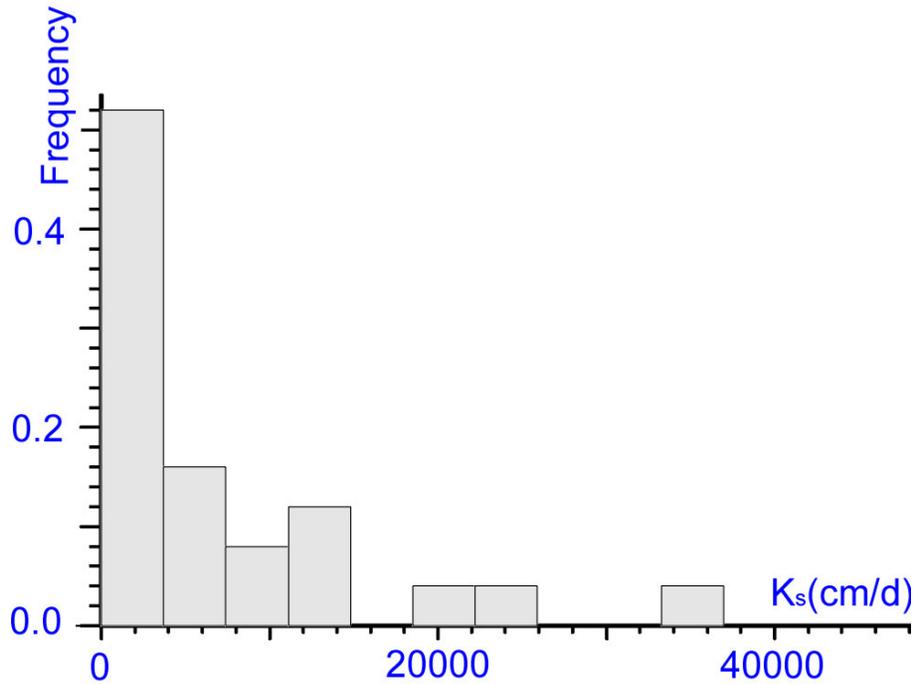
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**Fig. 6. (a)** Ring infiltrometer data histogram.

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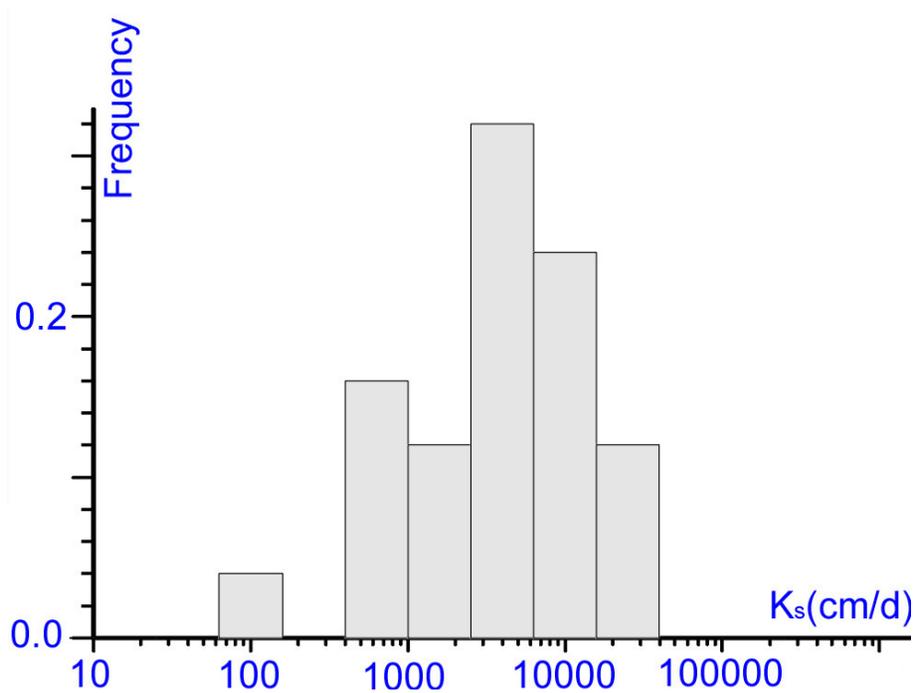


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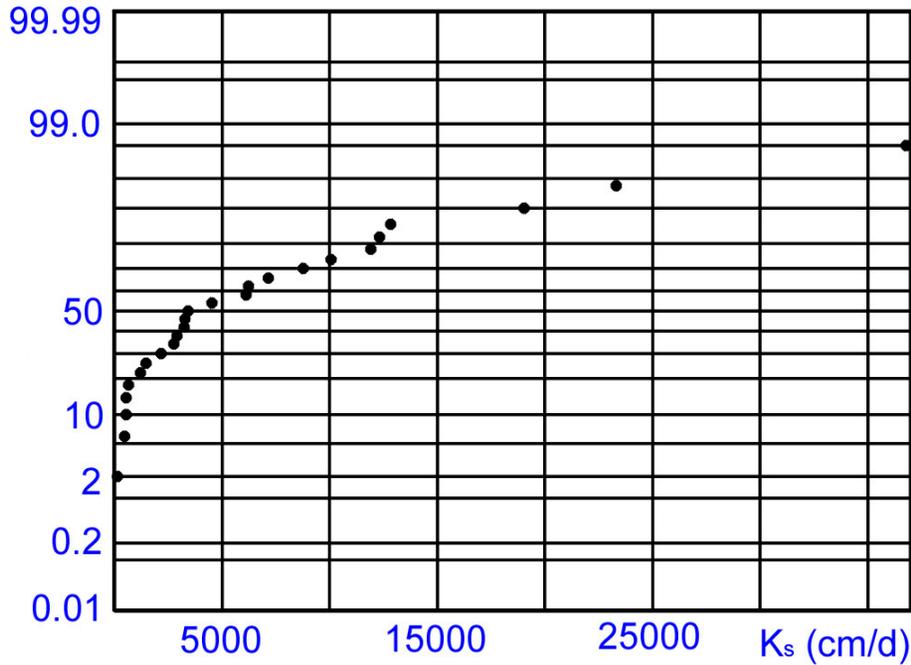


Fig. 6. (c) Ring infiltrometer data probability plot.

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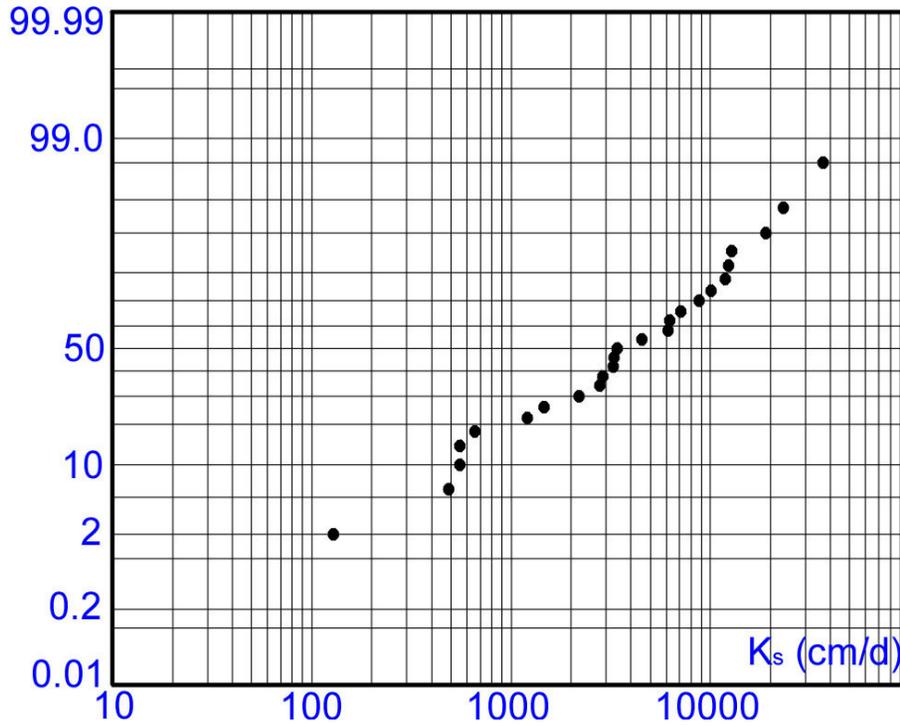
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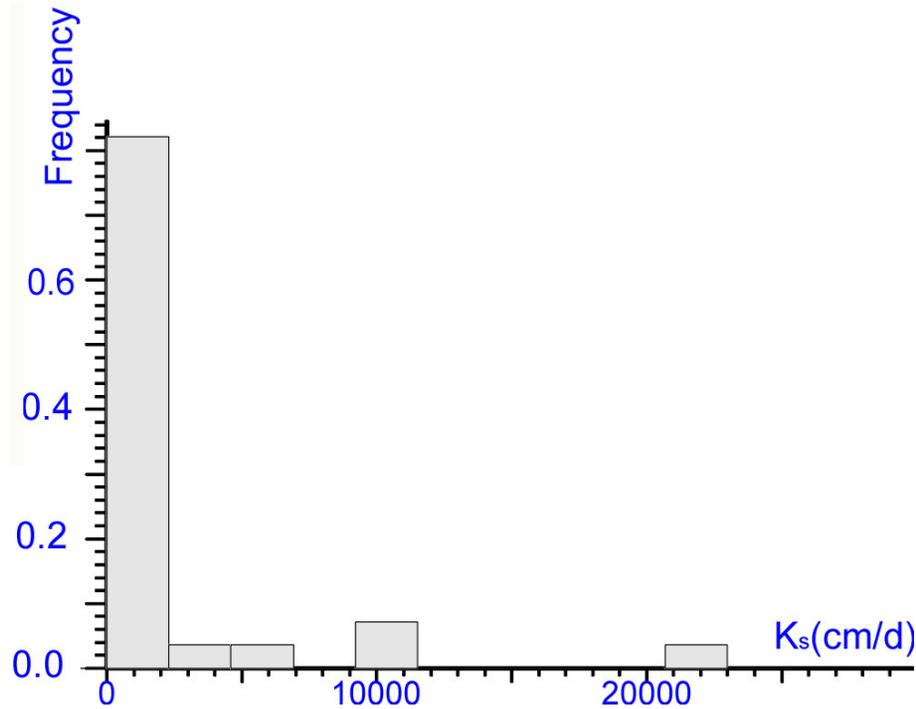


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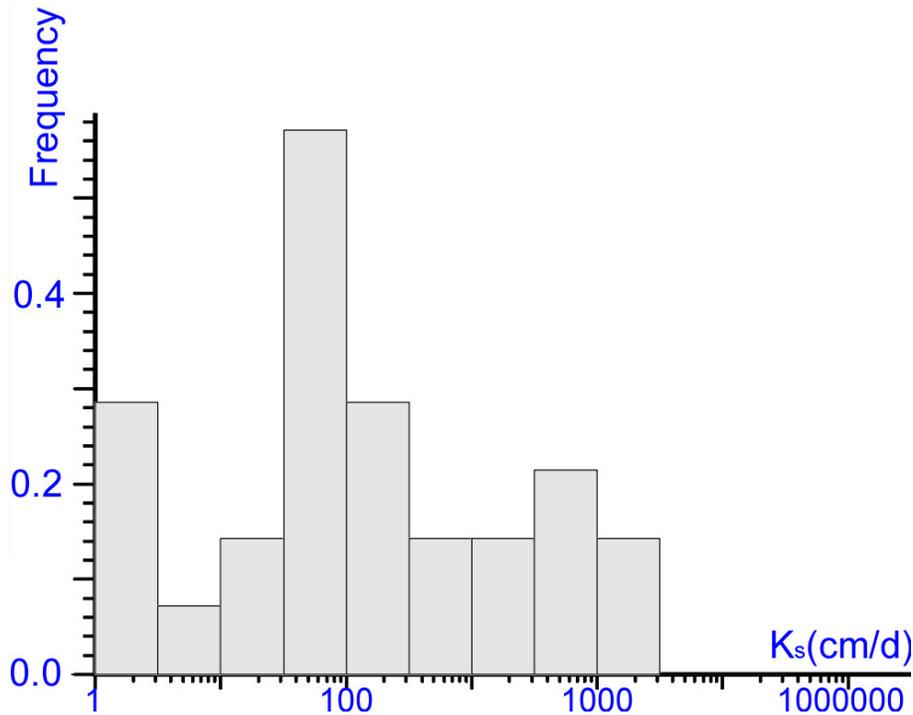


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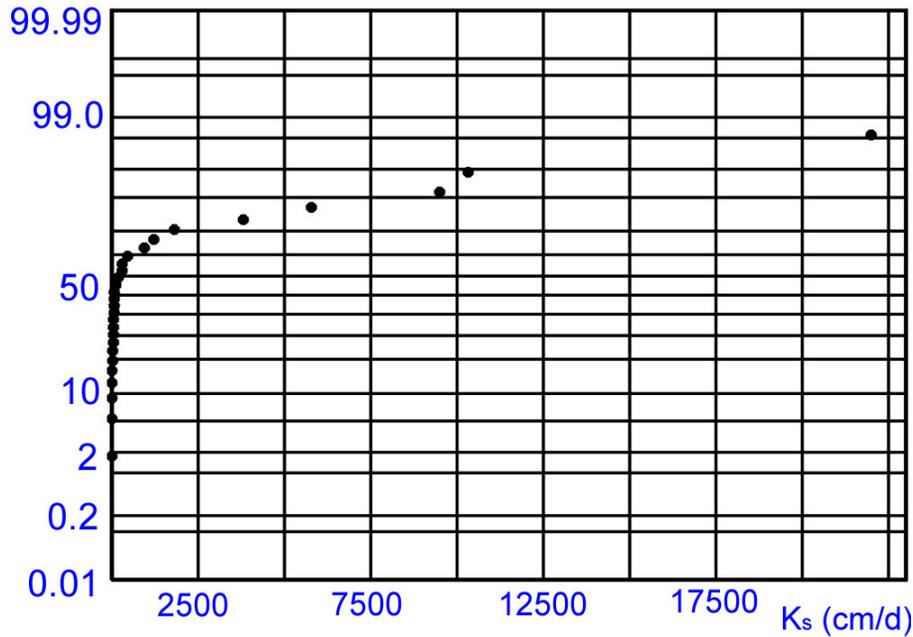


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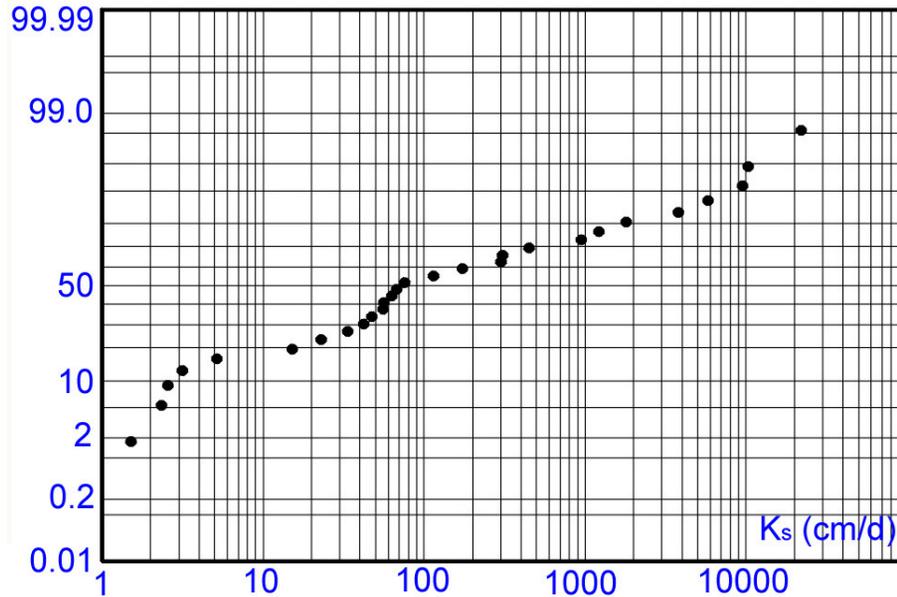


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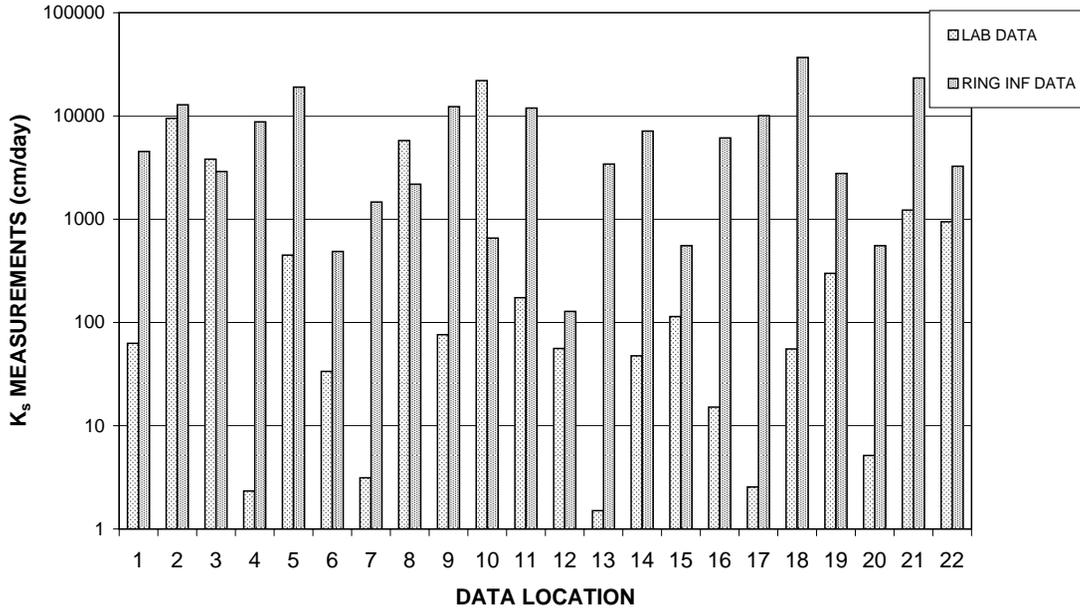
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**Fig. 8. (a)** Comparison between lab and ring infiltrometer  $K_s$ -data values.

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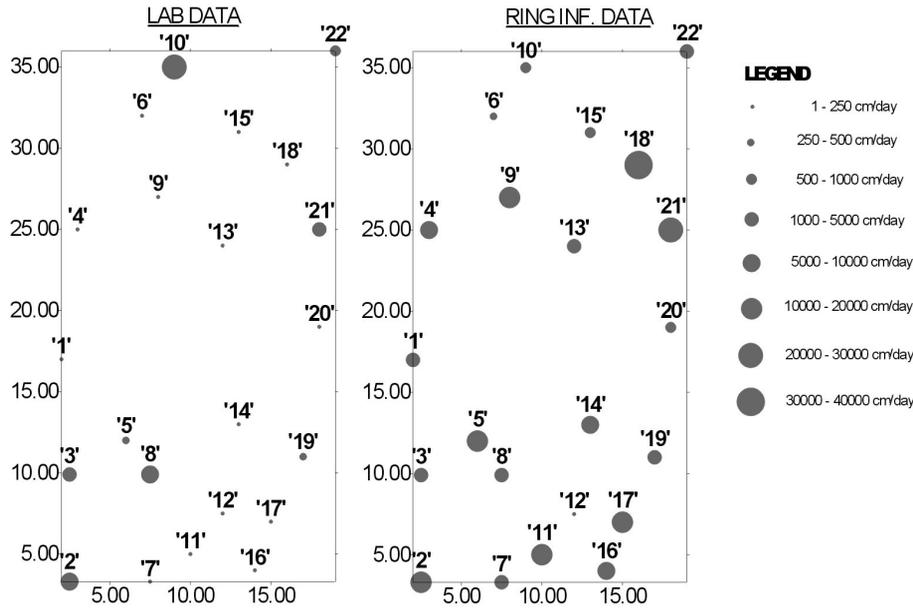


Fig. 8. (b) Classed posting of lab and ring infiltrometer  $K_s$ -data.



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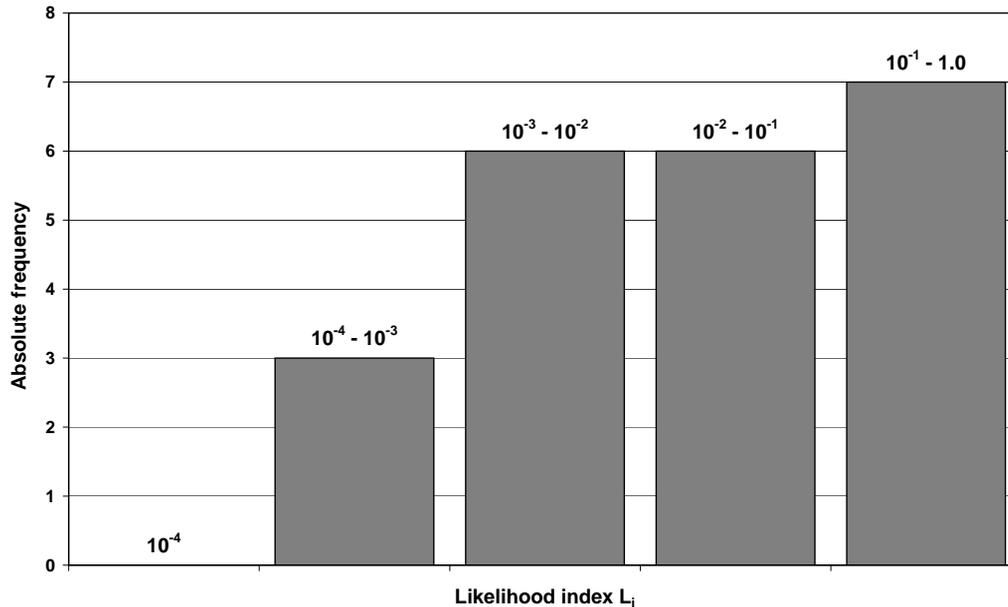


Fig. 9. (a) Absolute frequency histogram of likelihood index  $L_i$  for lab and ring infiltrometer data.

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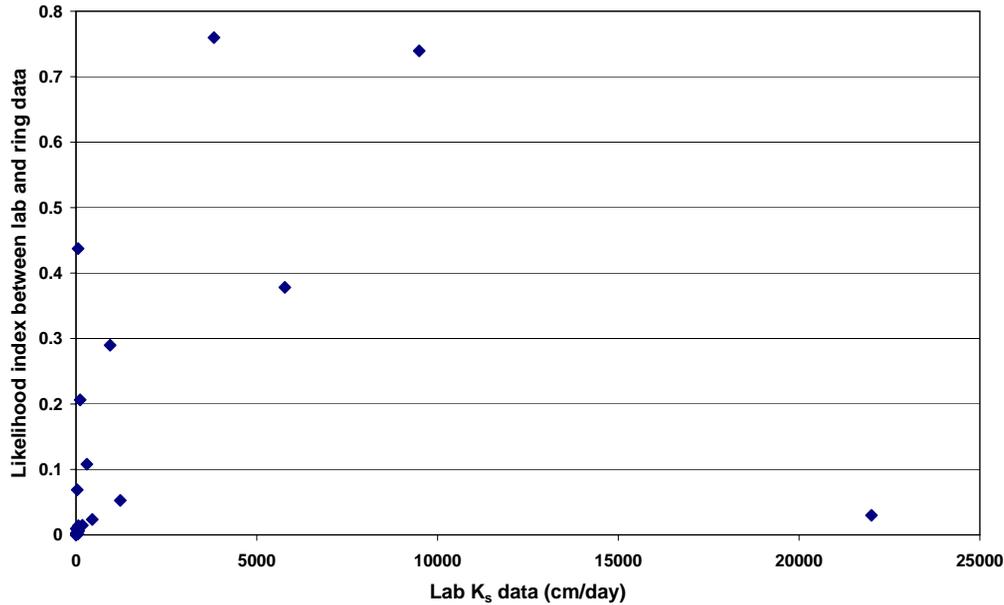


Fig. 9. (b) Likelihood index  $L_j$  vs. lab  $K_s$ -values.

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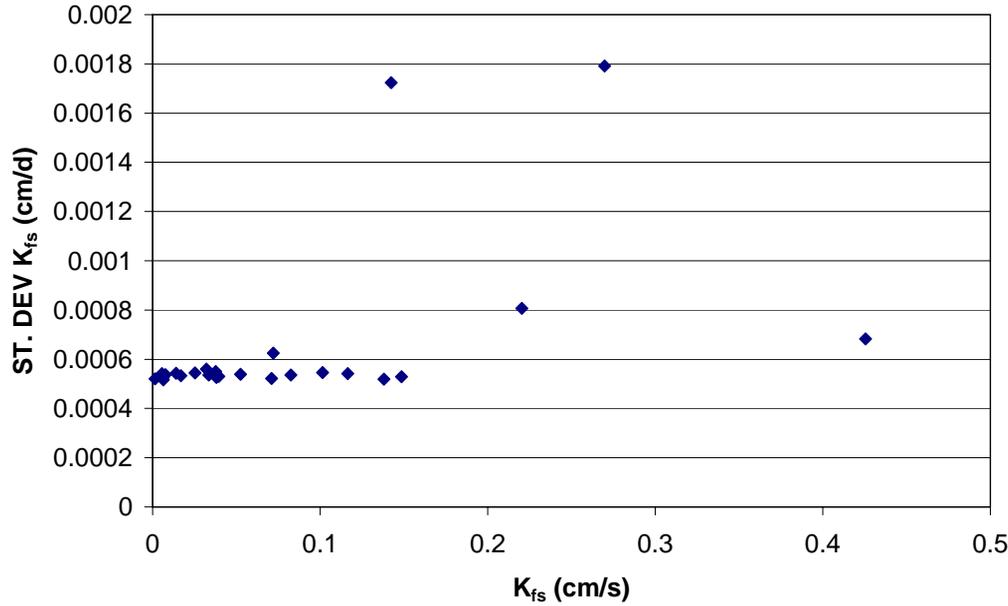


Fig. 10. (a) Ring infiltrometer data sensitivity analysis:  $K_{fs}$  standard deviation vs.  $K_{fs}$  mean.

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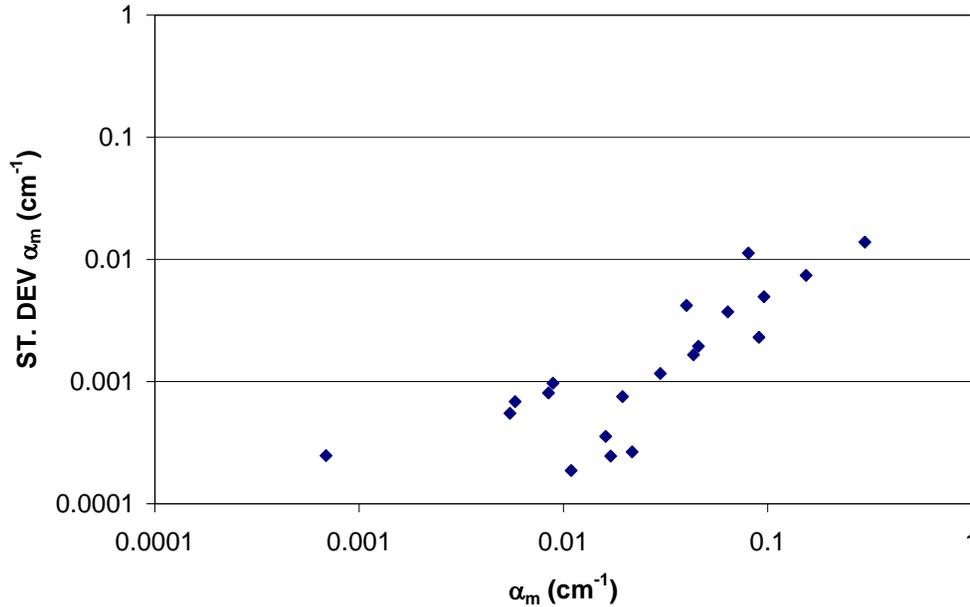
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**Fig. 10. (b)** Ring infiltrometer data sensitivity analysis:  $\alpha_m$  standard deviation vs.  $\alpha_m$  mean.

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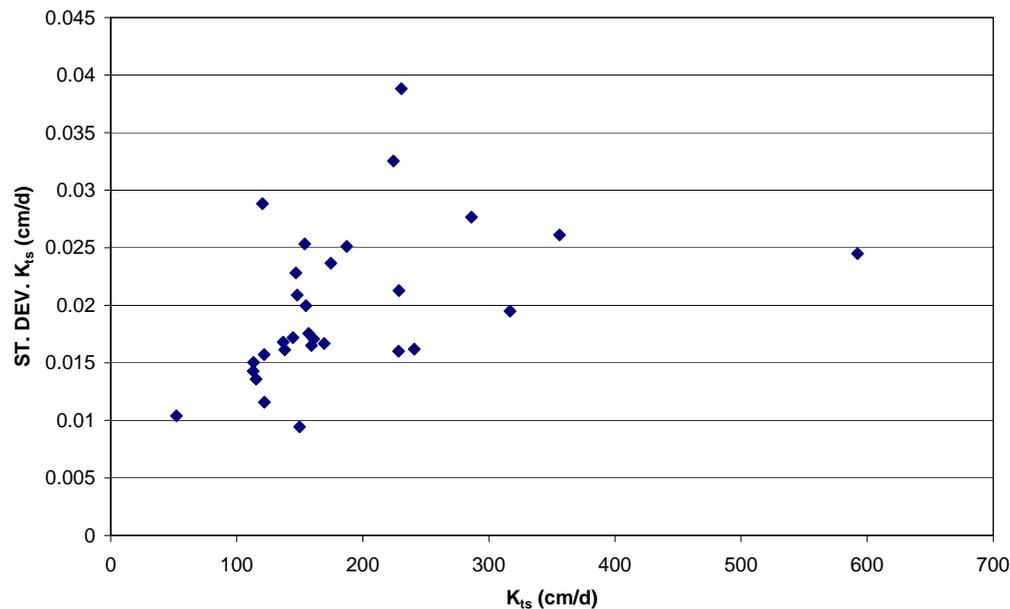


Fig. 11. (a) Disc infiltrometer data sensitivity analysis:  $K_{ts}$  standard deviation vs.  $K_{ts}$  mean values.

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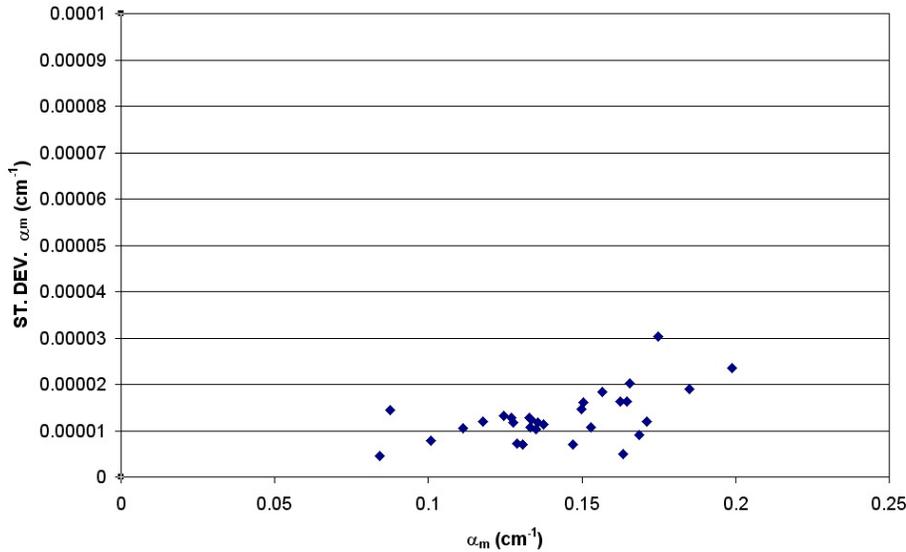


Fig. 11. (b) Disc infiltrometer data sensitivity analysis:  $\alpha_t$  standard deviation vs.  $\alpha_t$  mean values.

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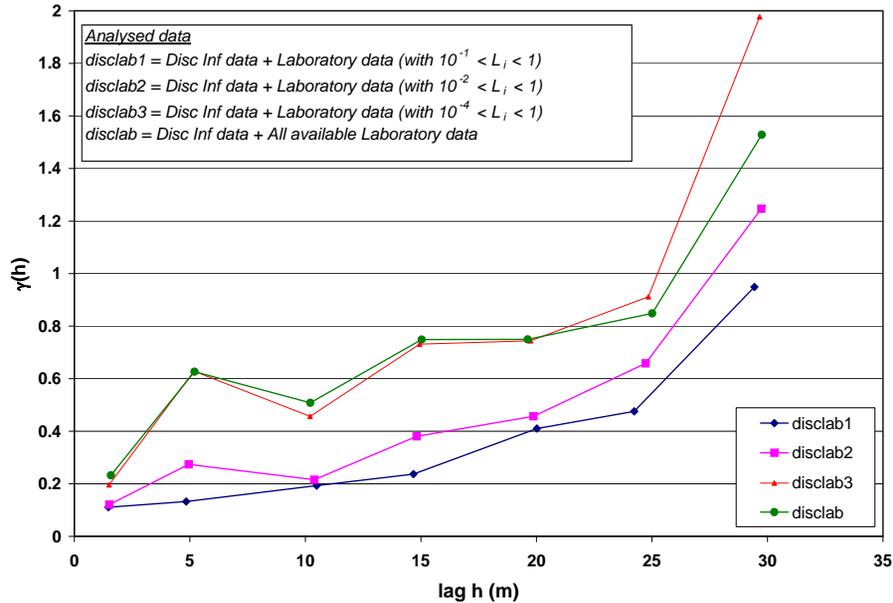
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**Fig. 12.** “Disclab” data sets experimental variograms.

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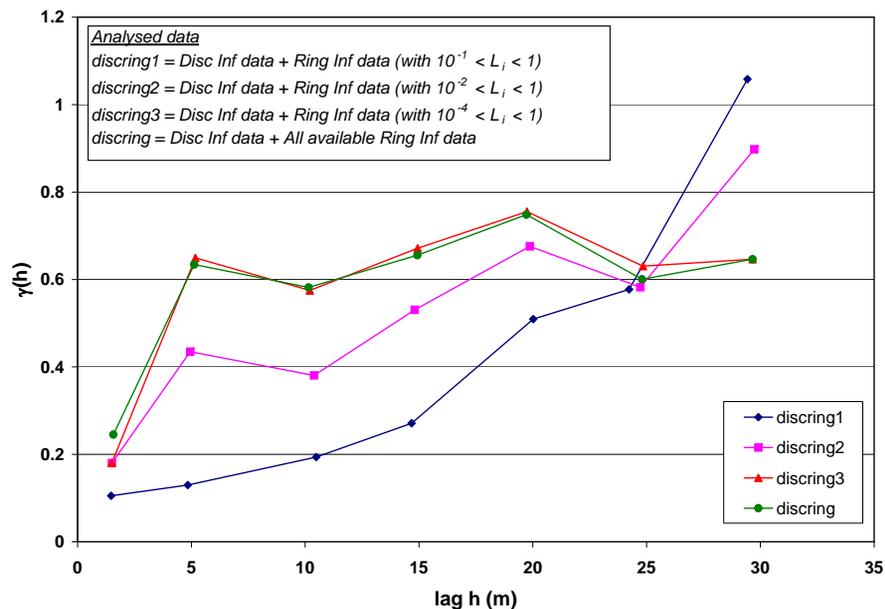


Fig. 13. “Discing” data sets experimental variograms.

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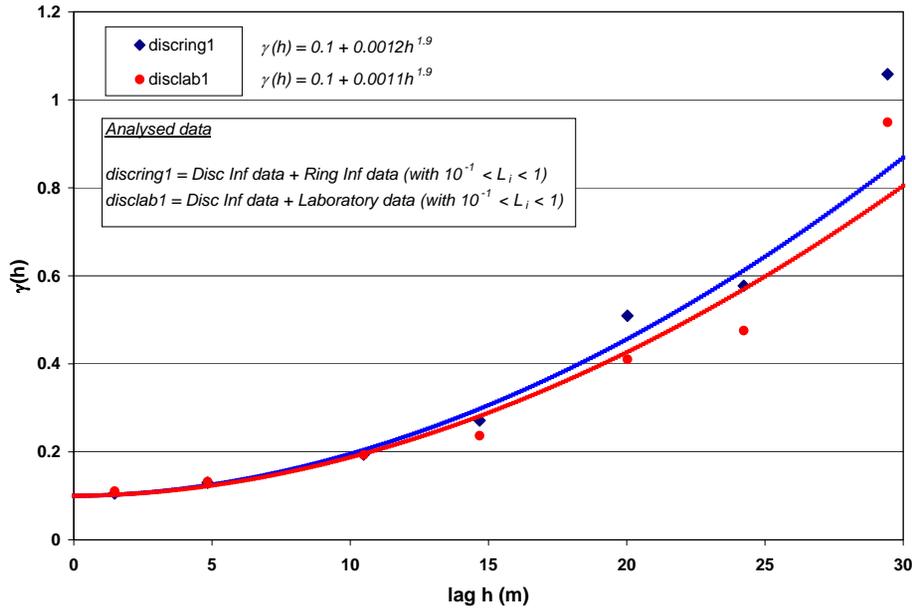


Fig. 14. (a) Variogram modelling: discing1 and disclab1 variogram models.

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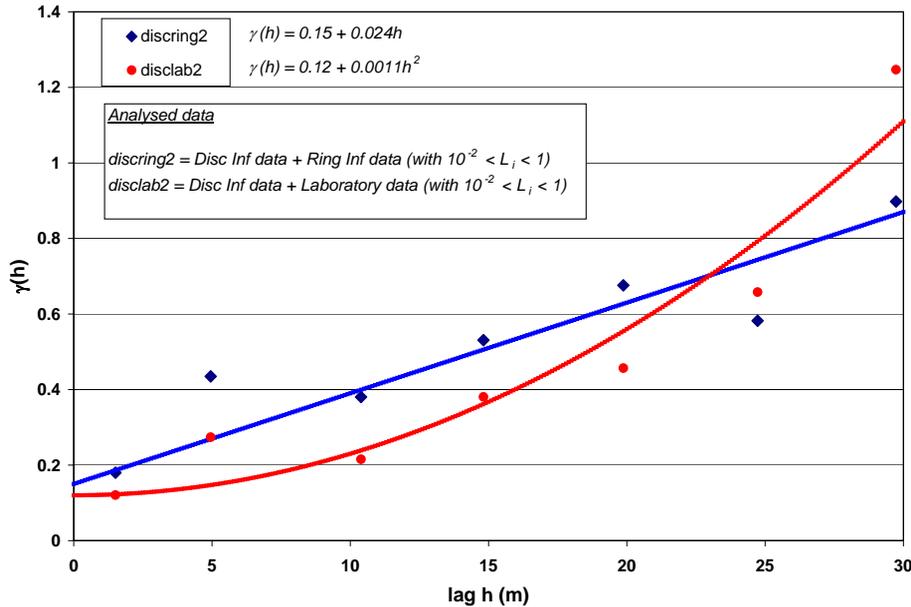


Fig. 14. (b) Variogram modelling: discing2 and disclab2 variogram models.

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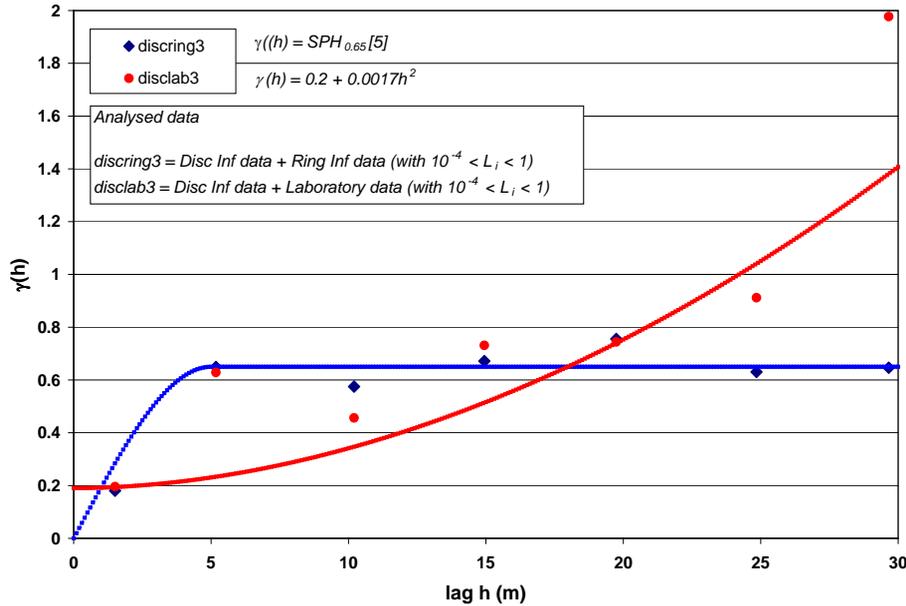


Fig. 15. (a) Variogram modelling: discing3 and disclab3 variogram models.

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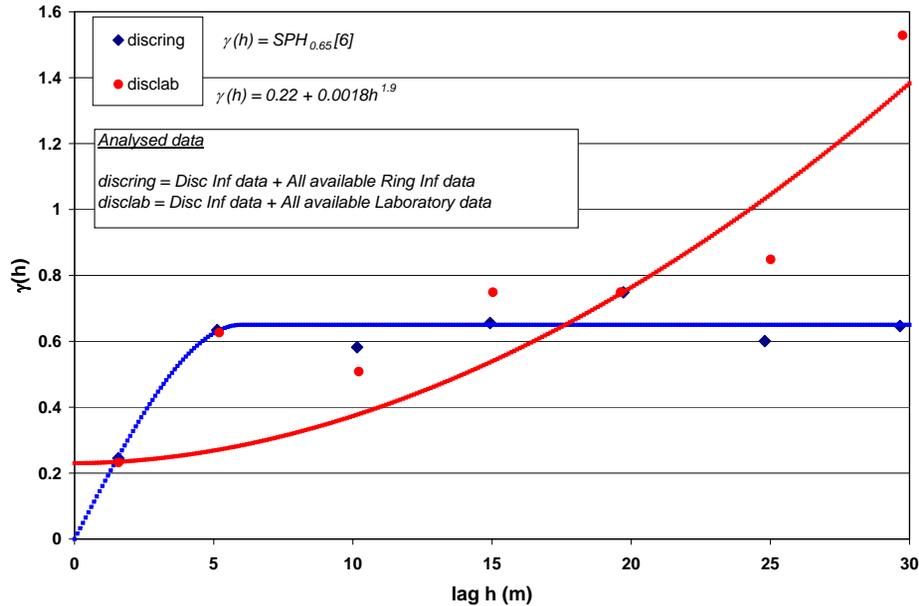


Fig. 15. (b) Variogram modelling: discing (all data) and disclab (all data) variogram models.

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