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Poroelasticity of a micro-heterogeneous material saturated by two immiscible fluids

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Abstract

The anisotropic poroelasticity framework presented by Aichi and Tokunaga (Int. J. Rock Mech. Min. Sci. 48(4) (2011) 580-584) for a porous material saturated by two immiscible fluids is generalized for micro-heterogeneous materials by introducing one additional poroelastic parameter. This framework is based on Coussy's thermodynamic framework, but is re-written using material parameters that are easier to evaluate experimentally. The presented generalization permits using the constitutive relations for modelling the behaviour of geomaterials with a heterogeneous solid phase composed of several constituents. The constitutive relations are given also for special cases of isotropic porous material saturated by two immiscible fluids or by one fluid.

Keywords: Poroelasticity, two phase fluid, constitutive relations, unsaturated, anisotropy

1. Introduction

The modelling of the mechanical behaviour of unsaturated porous materials is important for various applications in geotechnical engineering, petroleum geomechanics, biomechanics, etc. Based on the thermodynamics theory and Lagrangian saturation concept, Coussy [1, 2] presented a solid framework for modelling the elastoplastic behaviour of unsaturated porous media. This framework is presented based on material properties that are very difficult to evaluate experimentally. Aichi and Tokunaga [3] represented this theoretical framework for a more general case of anisotropic poroelasticity using material properties that are easier to evaluate experimentally. By means of a series of thought experiments, these authors established relationships for the unsaturated poroelasticity material parameters used in Coussy's framework as functions of porosity, saturated poroelasticity material parameters (drained stiffness and Biot effective stress coefficient tensors), Bishop effective stress coefficient and degree of saturation. But these relationships are limited to micro-homogeneous porous materials for which the solid phase is made up of one single constituent. It is well-known that this is not the case for many geomaterials constituted by a heterogeneous solid phase. The aim of this work is to generalize the presented constitutive relations in [3] by taking into account the heterogeneity of the solid phase of the porous material. This is done by introducing an additional poroelastic parameter into the formulations.

2. Poroelastic constitutive relations for unsaturated media

The poroelastic constitutive relations presented in the following are written considering an infinitesimal representative element of porous material. The material with total volume V is composed of a solid phase with volume V_s and a connected pore volume with volume V_ϕ . The occluded pore volume and the fluid filling it are considered to be a part of the solid phase. The solid phase itself may be composed of one or several constituents. The Lagrangian porosity is defined as $\phi = V_\phi/V_0$ and its variation is expressed by $\varphi = \phi - \phi_0$, where the subscript 0 is for referring to the initial

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21 state. It is assumed that the pore volume is composed of two parts, each one fully saturated by a fluid. These pore
 22 volumes are referred to in the following by subscripts wF and nwF, respectively for wetting fluid and non-wetting
 23 fluid, so that $\phi = \phi_{wF} + \phi_{nwF}$. From this relation one obtains

$$\varphi = \phi - \phi_0 = \varphi_{wF} + \varphi_{nwF} \quad (1)$$

24 The Lagrangian partial saturations for wetting and non-wetting fluids represent the fraction of the pore volume
 25 occupied by each fluid:

$$S_{wF} = \frac{\phi_{wF}}{\phi_0} \quad ; \quad S_{nwF} = \frac{\phi_{nwF}}{\phi_0} \quad ; \quad S_{wF} + S_{nwF} = 1 \quad (2)$$

26 The poroelastic constitutive relations for a general case of a fully anisotropic unsaturated material are given below
 27 [1, 2, 3]:

$$\sigma_{ij} = C_{ijkl}\varepsilon_{kl} - \alpha_{wFij}p_{wF} - \alpha_{nwFij}p_{nwF} \quad (3)$$

$$\varphi_{wF} = \alpha_{wFij}\varepsilon_{ij} + \frac{1}{N_{wF-wF}}p_{wF} + \frac{1}{N_{wF-nwF}}p_{nwF} \quad (4)$$

$$\varphi_{nwF} = \alpha_{nwFij}\varepsilon_{ij} + \frac{1}{N_{wF-nwF}}p_{wF} + \frac{1}{N_{nwF-nwF}}p_{nwF} \quad (5)$$

28 where C_{ijkl} is the stiffness tensor and α_{wFij} , α_{nwFij} , N_{wF-wF} , N_{wF-nwF} , $N_{nwF-nwF}$ are material constants that are
 29 equivalent to Biot effective stress coefficient tensor α_{ij} and Biot modulus N in saturated conditions. As mentioned
 30 by Aichi and Tokunga [3], these material parameters defined for unsaturated poroelasticity are not directly relevant
 31 to those which can be measured by laboratory experiments. In other words, an experimental evaluation of these pa-
 32 rameters is very difficult, if not impossible. To overcome this difficulty, Aichi and Tokunga [3] propose an interesting
 33 method, by means of a series of thought experiments, to establish relations among these unsaturated poroelasticity
 34 material properties with other parameters that are easier to evaluate experimentally. These relations which are pre-
 35 sented in [3] for the case of a micro-homogeneous porous material, are generalized in the following section for the
 36 case of a micro-heterogeneous porous material for which the solid phase is composed of several constituents. The
 37 derivation of these relations is done using the same method as in [3] and for this reason the structure of the following
 38 section is chosen similar to section (3) of [3].

39 3. Theoretical relationships among poroelastic coefficients

40 3.1. Uniform pore pressure build-up

41 When the pressure increments in wetting and non-wetting fluid are the same, equations 3 to 5 must be reduced
 42 to the saturated poroelasticity equations, $\sigma_{ij} = C'_{ijkl}\varepsilon_{kl} - \alpha_{ij}p$ and $\varphi = \alpha_{ij}\varepsilon_{ij} + p/N$ [1, 3]. In these equations C'_{ijkl}
 43 is the drained stiffness tensor, α_{ij} in Biot effective stress coefficient tensor and N is Biot modulus. By assuming
 44 $p_{wF} = p_{nwF} = p$ in equation 3 and comparing with $\sigma_{ij} = C'_{ijkl}\varepsilon_{kl} - \alpha_{ij}p$ one finds $C_{ijkl} = C'_{ijkl}$ and $\alpha_{wFij} + \alpha_{nwFij} = \alpha_{ij}$.
 45 This latter equation permits writing

$$\alpha_{wFij} = \chi\alpha_{ij} \quad ; \quad \alpha_{nwFij} = (1 - \chi)\alpha_{ij} \quad (6)$$

46 where χ is a constant depending on the Lagrangian saturation of wetting fluid. Replacing these relations in the
 47 constitutive relation 3 gives

$$\sigma_{ij} = C'_{ijkl}\varepsilon_{kl} - \alpha_{ij}[\chi p_{wF} + (1 - \chi)p_{nwF}] \quad (7)$$

48 which shows that the parameter χ is equivalent to Bishop [4] effective stress coefficient. Inserting equations 4 and
 49 5 and $p_{wF} = p_{nwF} = p$ in equation 1 and comparing with the saturated poroelasticity equation $\varphi = \alpha_{ij}\varepsilon_{ij} + p/N$, results
 50 in the following relation between the unsaturated poroelasticity constants and Biot modulus N for the saturated case
 51 [1]:

$$\frac{1}{N_{wF-wF}} + \frac{2}{N_{wF-nwF}} + \frac{1}{N_{nwF-nwF}} = \frac{1}{N} \quad (8)$$

3.2. Unjacketed condition

Theunjacketed condition is defined as $p_{wF} = p_{nwF} = p$ and $\sigma_{ij} = -p\delta_{ij}$. Under this condition, equal increments of confining pressure and pore pressure are simultaneously applied to the sample, as if the sample was submerged, without a jacket, into a fluid under the pressure p . Terzaghi effective stress $\sigma'_{ij} = \sigma_{ij} + p\delta_{ij}$ in this condition remains constant. In the unjacketed condition a micro-homogeneous porous material would deform as if all the pores were filled with the solid component. The skeleton and the solid component experience a uniform volumetric strain $dV/V_0 = dV_s/V_{s0} = dV_\phi/V_{\phi0}$, so that we can write [3]:

$$\varepsilon_{kk} = \frac{\varphi_{wF}}{\phi_0 S_{wF}} = \frac{\varphi_{nwF}}{\phi_0 S_{nwF}} \quad (9)$$

When the solid phase is heterogeneous the skeleton and the solid component will experience different volumetric strains and the above relation is modified as following:

$$\beta \varepsilon_{kk} = \frac{\varphi_{wF}}{\phi_0 S_{wF}} = \frac{\varphi_{nwF}}{\phi_0 S_{nwF}} \quad (10)$$

where the parameter β is a poroelastic parameter defined as the ratio of the volumetric strain of the pore volume to the one of the solid skeleton in unjacketed condition.

$$\beta = \left(\frac{dV_\phi/V_{\phi0}}{dV/V_0} \right)_{\substack{p_{wF}=p_{nwF}=p \\ \sigma_{ij}=-p\delta_{ij}}} \quad (11)$$

In poroelasticity theory, two elastic moduli K_s and K_ϕ are classically defined for unjacketed condition [5, 6, 7].

$$\frac{1}{K_s} = -\frac{1}{V_0} \left(\frac{\partial V}{\partial p} \right)_{\substack{p_{wF}=p_{nwF}=p \\ \sigma_{ij}=-p\delta_{ij}}} ; \quad \frac{1}{K_\phi} = -\frac{1}{V_{\phi0}} \left(\frac{\partial V_\phi}{\partial p} \right)_{\substack{p_{wF}=p_{nwF}=p \\ \sigma_{ij}=-p\delta_{ij}}} \quad (12)$$

The poroelastic parameter β can thus be identified as:

$$\beta = \frac{K_s}{K_\phi} \quad (13)$$

The parameters K_s and K_ϕ are equivalent respectively to κ_M and κ_ϕ defined by Brown and Korringa [5] and K'_s and K''_s defined by Rice and Cleary [6]. For a micro-homogeneous porous material $K_s = K_\phi = K_m$ where K_m is the bulk modulus of the single solid constituent of the porous material. In the case of a porous material which is composed of two or more solids and therefore is heterogeneous at the micro-scale, the unjacketed modulus K_s is some weighted average of the bulk moduli of solid constituents [7]. What this average should be is generally unknown, however, Ghabezloo and Sulem [8] evaluated the unjacketed modulus of Rothbach sandstone using Hill's [9] average formula and found a good accordance with the experimentally evaluated modulus. The modulus K_ϕ for such a material has a complicated dependence on the material properties. Generally it is not bounded by the elastic moduli of the solid components and can even have a negative sign if the bulk moduli of the individual solid components are greatly different one from another [10]. The modulus K_s can be easily evaluated in an unjacketed test by measurement of the variations of the total volume of the tested sample. The variation of the pore volume of the sample in this test, evaluated from the quantity of fluids exchanged between the sample and the pore pressure generators could in principle give the modulus K_ϕ . However experimental evaluation of this parameter is very difficult as the volume of the exchanged fluid has to be corrected for the effect of fluid compressibility, and also for the effect of the deformations of the pore pressure generator and drainage system in order to access to the variation of the pore volume of the sample [11, 12]. However this parameter can be evaluated indirectly as a function of the other poroelastic parameters which can be measured independently, as presented in [11, 13]. It is also possible to evaluate this parameter using the homogenization method [14, 15]. The parameter β for a hardened cement paste is evaluated equal to 1.24 from the analysis of poromechanics tests [14] ($K_s = 21.0$ GPa, $K_\phi = 16.9$ GPa) and equal to 1.1 using the homogenization

84 method [11]. From the experimental results of Hart and Wang [13] the parameter β can be evaluated equal to 6.6 for
 85 Berea sandstone ($K_s = 28.9$ GPa, $K_\phi = 4.4$ GPa) and 9.7 for Indiana limestone ($K_s = 72.6$ GPa, $K_\phi = 7.5$ GPa).
 86 Based on the experimental results of Laurent et al. [16] cited in [13], the parameter β can be evaluated between 0.33
 87 and 0.56 for Vilhonneur limestone ($K_s = 41$ and 56 GPa from two independent measurements, K_ϕ between 100 and
 88 125 GPa).

89 By re-writing equations 3, 4 and 5 for the unjacketed condition ($p_{wF} = p_{nwF} = p$ and $\sigma_{ij} = -p\delta_{ij}$) and using
 90 equations 10 and 6 the following relations are found between the Biot moduli for unsaturated condition and the other
 91 poroelastic parameters:

$$\frac{1}{N_{wF-wF}} + \frac{1}{N_{wF-nwF}} = (\alpha_{ij}\chi - \phi_0\beta S_{wF}\delta_{ij}) C'_{ijkl}{}^{-1} (\delta_{kl} - \alpha_{kl}) \quad (14)$$

$$\frac{1}{N_{wF-wF}} + \frac{1}{N_{nwF-nwF}} = [\alpha_{ij}(1 - \chi) - \phi_0\beta(1 - S_{wF})\delta_{ij}] C'_{ijkl}{}^{-1} (\delta_{kl} - \alpha_{kl}) \quad (15)$$

92 3.3. Build-up of the pressure of the wetting phase fluid under constant Lagrangian partial porosity of the non-wetting 93 phase fluid and no deformation of the solid constituent

94 In this thought experiment the build-up of the pressure of the wetting phase fluid under constant Lagrangian partial
 95 porosity of the non-wetting phase fluid are considered as follows:

$$p_{wF} = p; \quad \varphi_{nwF} = 0; \quad S_{wF} = \text{const.} \quad (16)$$

96 Schematic representations of these conditions are presented in [3]. In addition to these conditions, the total stress
 97 is supposed to be:

$$\sigma_{ij} = -\beta\phi_0 [S_{wF}p + (1 - S_{nwF})p_{nwF}] \delta_{ij} \quad (17)$$

98 By assuming $\beta = 1$ the above equation is reduced to the one presented in [3]. Using equation (17), following the
 99 same way as presented in section 3.3 of [3], one can show that $\varphi_{wF} + \varphi_{nwF} = \varepsilon_{kk}$ and the strain of the solid constituent
 100 is zero. Finally, the following relations are obtained for the three Biot moduli (see [3] for more details on derivation):

$$\frac{1}{N_{wF-nwF}} = (\alpha_{ij}\chi - \phi_0\beta S_{wF}\delta_{ij}) C'_{ijkl}{}^{-1} (\delta_{kl} - \alpha_{kl}) \quad (18)$$

$$\times [\alpha_{mn}(1 - \chi) - \phi_0\beta(1 - S_{wF})\delta_{mn}] (\alpha_{mn} - \phi_0\beta\delta_{mn})^{-1}$$

$$\frac{1}{N_{wF-wF}} = (\alpha_{ij}\chi - \phi_0\beta S_{wF}\delta_{ij}) C'_{ijkl}{}^{-1} (\delta_{kl} - \alpha_{kl}) \quad (19)$$

$$\times (\alpha_{mn}\chi - \phi_0\beta S_{wF}\delta_{mn}) (\alpha_{mn} - \phi_0\beta\delta_{mn})^{-1}$$

$$\frac{1}{N_{nwF-nwF}} = [\alpha_{ij}(1 - \chi) - \phi_0\beta(1 - S_{wF})\delta_{ij}] C'_{ijkl}{}^{-1} (\delta_{kl} - \alpha_{kl}) \quad (20)$$

$$\times [\alpha_{mn}(1 - \chi) - \phi_0\beta(1 - S_{wF})\delta_{mn}] (\alpha_{mn} - \phi_0\beta\delta_{mn})^{-1}$$

101 Equations 6, 18, 19 and 20 give relationships for unsaturated poroelasticity material parameters as functions of
 102 other parameters which are easier to evaluate experimentally.

103 4. Poroelastic constitutive relations for micro-heterogeneous materials

104 4.1. General case of anisotropic material

105 Substituting equations 6, 18, 19 and 20 in the constitutive relations 3, 4 and 5 permit writing the poroelastic
106 constitutive relations for an anisotropic micro-heterogeneous material:

$$\sigma_{ij} = C'_{ijkl} \varepsilon_{kl} - \alpha_{ij} [\chi p_{wF} + (1 - \chi) p_{nwF}] \quad (21)$$

$$\begin{aligned} \varphi_{wF} = & \alpha_{ij} \chi \varepsilon_{ij} + (\alpha_{ij} \chi - \phi_0 \beta S_{wF} \delta_{ij}) C'^{-1}_{ijkl} (\delta_{kl} - \alpha_{kl}) (\alpha_{mn} \chi - \phi_0 \beta S_{wF} \delta_{mn}) \\ & \times (\alpha_{mn} - \phi_0 \beta \delta_{mn})^{-1} p_{wF} + (\alpha_{ij} \chi - \phi_0 \beta S_{wF} \delta_{ij}) C'^{-1}_{ijkl} (\delta_{kl} - \alpha_{kl}) \\ & \times [\alpha_{mn} (1 - \chi) - \phi_0 \beta (1 - S_{wF}) \delta_{mn}] (\alpha_{mn} - \phi_0 \beta \delta_{mn})^{-1} p_{nwF} \end{aligned} \quad (22)$$

$$\begin{aligned} \varphi_{nwF} = & \alpha_{ij} (1 - \chi) \varepsilon_{ij} + (\alpha_{ij} \chi - \phi_0 \beta S_{wF} \delta_{ij}) C'^{-1}_{ijkl} (\delta_{kl} - \alpha_{kl}) \\ & \times [\alpha_{mn} (1 - \chi) - \phi_0 \beta (1 - S_{wF}) \delta_{mn}] (\alpha_{mn} - \phi_0 \beta \delta_{mn})^{-1} p_{wF} \\ & + (\alpha_{ij} (1 - \chi) - \phi_0 \beta (1 - S_{wF}) \delta_{ij}) C'^{-1}_{ijkl} (\delta_{kl} - \alpha_{kl}) \\ & \times [\alpha_{mn} (1 - \chi) - \phi_0 \beta (1 - S_{wF}) \delta_{mn}] (\alpha_{mn} - \phi_0 \beta \delta_{mn})^{-1} p_{nwF} \end{aligned} \quad (23)$$

107 By assuming $\beta = 1$ the above equations are reduced to the ones presented in [3].

108 4.2. Special case of isotropy

109 In the case of isotropy the constitutive relations are reduced to:

$$\sigma_{ij} + \alpha [\chi p_{wF} + (1 - \chi) p_{nwF}] = 2G \varepsilon_{ij} + \left(K - \frac{2G}{3} \right) \varepsilon_{kk} \delta_{ij} \quad (24)$$

$$\begin{aligned} \varphi_{wF} = & \alpha \chi \varepsilon_{kk} \\ & + \frac{(\alpha \chi - \phi_0 \beta S_{wF})^2 (1 - \alpha)}{K (\alpha - \phi_0 \beta)} p_{wF} \\ & + \frac{(\alpha \chi - \phi_0 \beta S_{wF}) [\alpha (1 - \chi) - \phi_0 \beta (1 - S_{wF})] (1 - \alpha)}{K (\alpha - \phi_0 \beta)} p_{nwF} \end{aligned} \quad (25)$$

$$\begin{aligned} \varphi_{nwF} = & \alpha (1 - \chi) \varepsilon_{kk} \\ & + \frac{(\alpha \chi - \phi_0 \beta S_{wF}) [\alpha (1 - \chi) - \phi_0 \beta (1 - S_{wF})] (1 - \alpha)}{K (\alpha - \phi_0 \beta)} p_{wF} \\ & + \frac{[\alpha (1 - \chi) - \phi_0 \beta (1 - S_{wF})]^2 (1 - \alpha)}{K (\alpha - \phi_0 \beta)} p_{nwF} \end{aligned} \quad (26)$$

110 4.3. Special case of isotropy and saturation by one fluid

111 For an isotropic porous material saturated by one fluid we have $\chi = 1$ and $S_{wF} = 1$. The constitutive relations in
112 this case are reduced to:

$$\sigma_{ij} + \alpha p = 2G \varepsilon_{ij} + \left(K - \frac{2G}{3} \right) \varepsilon_{kk} \delta_{ij} \quad (27)$$

$$\varphi = \alpha \varepsilon_{kk} + \frac{(\alpha - \phi_0 \beta) (1 - \alpha)}{K} p \quad (28)$$

113 The term $(\alpha - \phi_0\beta)(1 - \alpha)/K$ in equation (28) is equal to $1/N$ where N is Biot modulus. Replacing $\alpha = 1 - K/K_s$
114 and $\beta = K_s/K_\phi$ in this relation the expression of Biot modulus presented in [17, 14] is retrieved:

$$\frac{1}{N} = \frac{\alpha}{K_s} - \frac{\phi_0}{K_\phi} \quad (29)$$

115 In the case of a micro-homogeneous porous material composed of one single solid constituent we have $K_s = K_\phi$
116 and equation 29 is reduced to the well-known relation $1/N = (\alpha - \phi_0)/K_s$ [17, 18].

117 5. Conclusions

118 The anisotropic poroelasticity framework presented by Aichi and Tokunaga [3] for a porous material saturated by
119 two immiscible fluids is generalized by taking into account the heterogeneity of the solid phase of the material. This
120 is done by introducing one additional poroelastic parameter β into the formulations. The parameter β is defined as
121 the ratio of the volumetric strain of the pore volume to the one of the solid skeleton inunjacketed condition. The
122 assumption of a homogeneous solid phase is introduced in [3] and also in [1] to obtain the relationships between
123 Biot modulus N and the other poroelastic parameters. The presented generalization permits using the unsaturated
124 poroelasticity constitutive relations for modelling the behaviour of anisotropic geomaterials with a heterogeneous
125 solid phase. The constitutive relations are presented also for special cases of isotropic porous material saturated by
126 two immiscible fluids as well as isotropic porous material saturated by one fluid.

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